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# Case study: measurement of Q and cumulative attenuation from VSP data

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NYSE: **DVN** devonenergy.com

#### Downgoing wave





#### Extracted analysis ribbon A 300ms window flattened on the first break.





Analysis ribbon



Amp: independent

#### Extracted analysis ribbon, spatially averaged with delz=400

After a spatial mix. Essentially each trace is the average of neighboring traces over +/- 400m







#### Analysis ribbon after spatial mix

## f-x amplitude spectrum of analysis ribbon



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#### f-x amplitude spectrum of analysis ribbon, delz=400

After a spatial mix. Essentially each trace is the average of neighboring traces over +/- 400m







#### Analysis ribbon after spatial mix: f-x amplitude spectrum

#### Spectral-ratio method of Q estimation This is the most commonly used method



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Predicted values for Q and T depend on the frequency range of the fit.

#### Dominant frequency method of Q estimation This is new. It has the virtue of not needing a division.





Strategy: Find the forward Q filter that when applied to  $w_1$  reduces the dominant frequency to that of  $w_2$ .



Related to Quan and Harris, 1997, but significantly different.

## Spectral-ratio analysis plots



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## Dominant-frequency analysis plots



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# Cumulative attenuation or CA after Hauge (1981)

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Attenuation  $(Q^{-1})$  consists of intrinsic and stratigraphic parts which combine to give total attenuation:

$$CA = \frac{\pi(t_2 - t_1)}{Q_{eff}} = \frac{\pi(t_2 - t_1)}{Q_{intrinsic}} + \frac{\pi(t_2 - t_1)}{Q_{strat}} \}$$

1. Effective attenuation is what is always measured.

2. Intrinsic attenuation is a rock or reservoir property. Monotonic.

3. Stratigraphic attenuation is an interference effect from short-path multiples (O'Doherty and Anstey 1971).



With time in the numerator,

this is called *cumulative* 

attenuation or CA.

#### *Q* versus *CA A test on a noise-free synthetic*





------ SRM (spectral ratio) —— DFM (dominant frequency) **CREWES** 

#### Q and CA estimates on raw data Reference level 2185 ft









## Q and CA estimates on spatially averaged data delz=100 *Reference level 2185 ft*



## Q and CA estimates on spatially averaged data delz=200 *Reference level 2185 ft*



## Q and CA estimates on spatially averaged data delz=300 *Reference level 2185 ft*



## Q and CA estimates on spatially averaged data delz=400 *Reference level 2185 ft*



## Q and CA estimates on spatially averaged data delz=500 *Reference level 2185 ft*



## Extending the estimates to the surface

Going from reference level 2185 ft. to reference level 0 ft.



- Attenuation measurements require the comparison of two signals.
- Measurements thus far compare the "shallowest" receiver to deeper receivers. Therefore all attenuation measures are relative to the depth of the shallowest receiver.
- Extending measurements to z = 0 requires knowledge of the signal at that depth.
- We have no receiver there, but, in theory we know the amplitude spectrum of the source wavelet.
- Two source options:
  - 1) Source wavelet is Klauder wavelet
  - 2) Source wavelet is time-derivative of Klauder wavelet

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#### The wavelets





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#### Modified analysis ribbon, with Klauder wavelet in column 1

After a spatial mix. Essentially each trace is the average of neighboring traces over +/- 500m



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Analysis ribbon with Klauder wavelet

#### f-x amplitude spectrum, with Klauder wavelet in column 1







## Q and CA estimates on spatially averaged data delz=500 **CREWES**

Reference level 0 ft, Source wavelet: Klauder 8-96Hz





#### Q and CA estimates on spatially averaged data delz=500 *Reference level 0 ft, Source wavelet: Time derivative Klauder 8-96Hz*



#### Conclusions





- Measurement of attenuation can be considered either as *Q* or as *CA*. The latter is more stable when attenuation is low.
- Two measurement algorithms were presented:
  - 1) Spectral-ratio method: very sensitive to amplitude balancing and frequency range
  - 2) Dominant-frequency method: insensitive to amplitudes and less sensitive to frequency range.
- When applied to real VSP data, very small residual upgoing waves cause instability.
- Both methods gave similar results on spatially averaged data.
- Extension to the surface assuming a known source was investigated.
- Overall average Q values seem quite low when referenced to the surface.
  - Relative to 2185 ft, Q ranged from 60 to 110.
  - Relative to 0 ft, Q ranged from less than 20 to 55.
- This study was conducted with software in the CREWES Matlab toolbox.





I think the sponsors of CREWES, especially Devon Canada, for their support.

- Devon USA made the VSP data available.
- Colleagues at Devon provided valuable commentary and insight.

## Spectral-ratio analysis plots



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