

Inversion for indicators of interconnected / aligned cracks

Joint PP and PS AVAZ inversion for orthogonal fracture weaknesses

Huaizhen Chen and Kris Innanen

Calgary, AB, Canada

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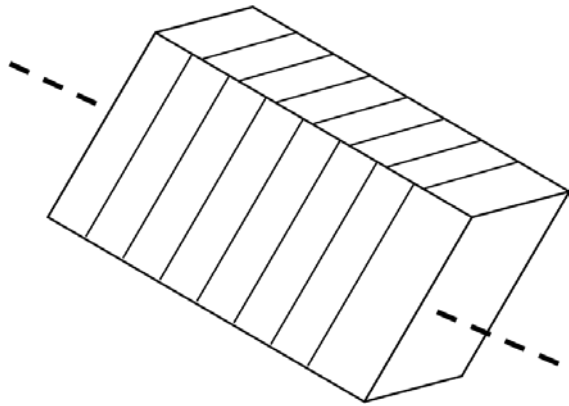
- Introduction
- Theory and method
- Numerical examples
- Conclusions



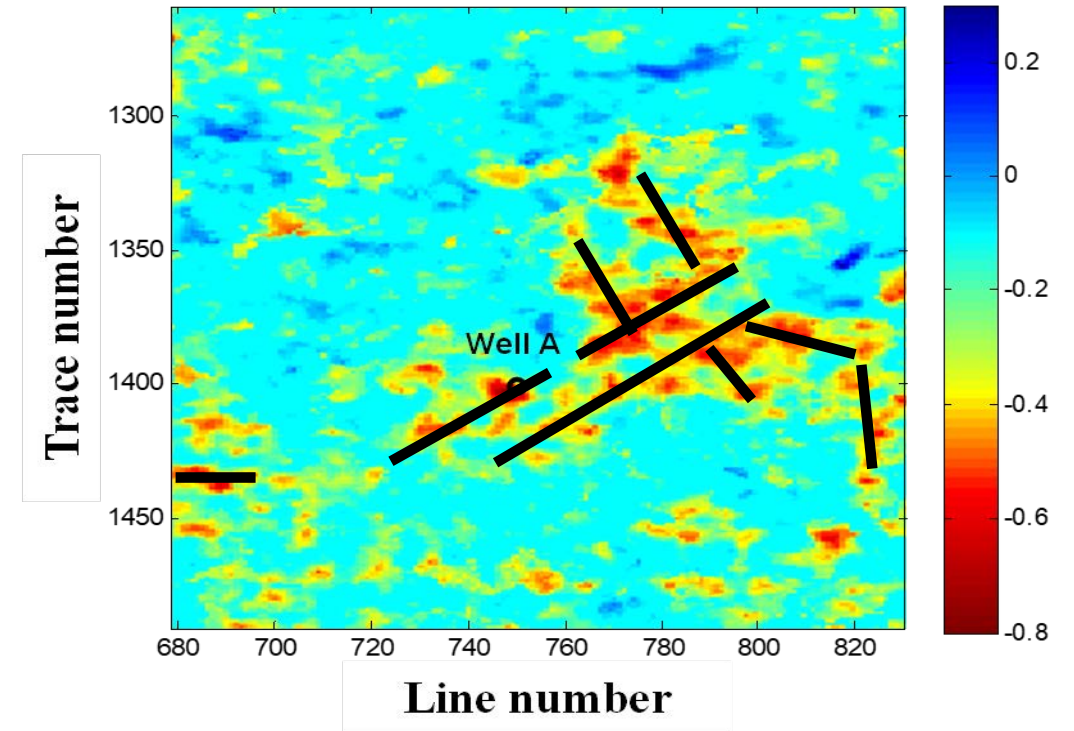
- Assumption of single set of fractures



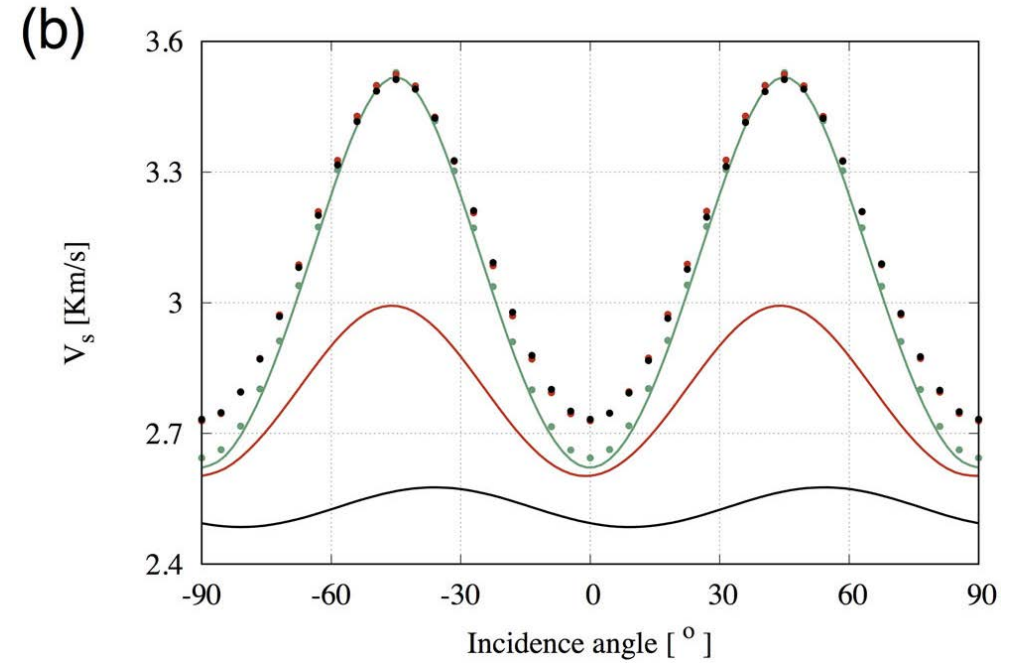
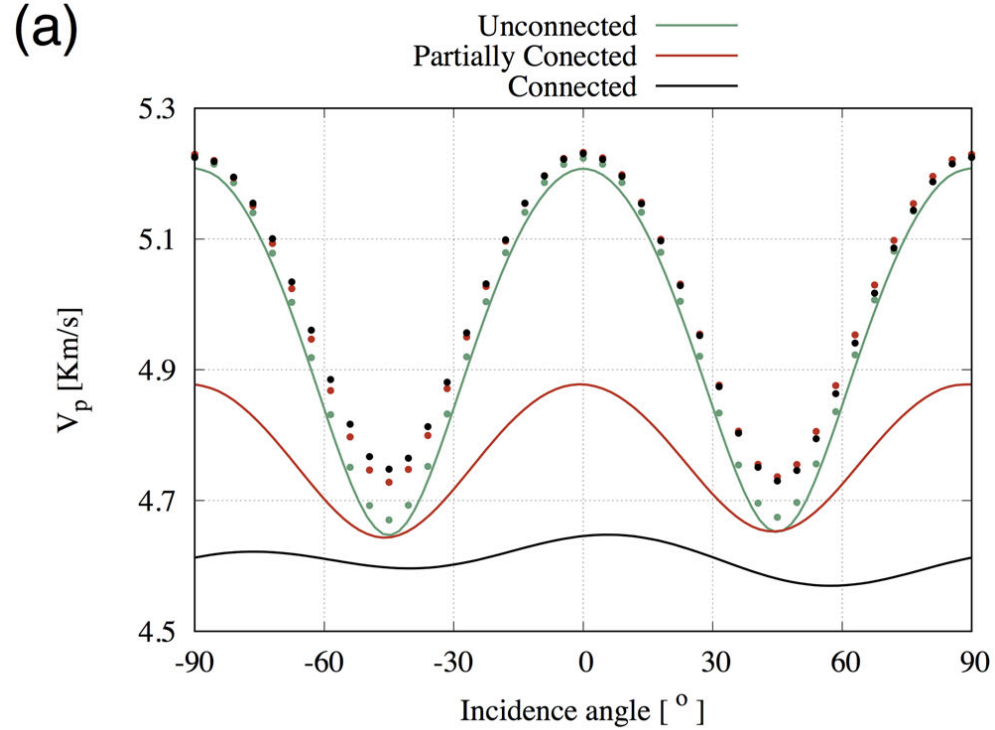
HTI media



TTI media



Fracture connectivity?



Fracture connectivity can reduce the velocity anisotropy of seismic waves.
(Rubino et al., GJI, 2017, 210, 223–227)

How is reflection coefficient influenced by connected fractures?



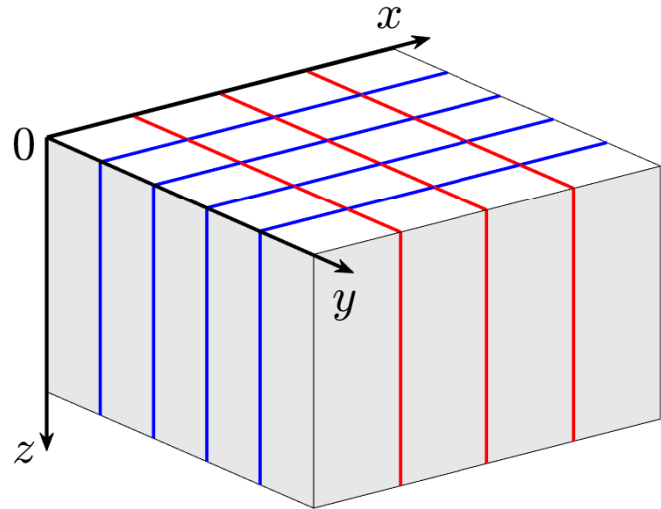
Objective

- PP- and PS-wave reflection coefficients in rocks containing interconnected fractures
- Azimuthal AVO inversion for interconnected fracture indicators



Theory and method: PP- and PS-wave reflection coefficients

An effective model of two orthogonal fractures in an isotropic background



$$\mathbf{C} = \begin{bmatrix} C_{11} & C_{12} & C_{13} & 0 & 0 & 0 \\ C_{12} & C_{22} & C_{23} & 0 & 0 & 0 \\ C_{13} & C_{23} & C_{33} & 0 & 0 & 0 \\ 0 & 0 & 0 & C_{44} & 0 & 0 \\ 0 & 0 & 0 & 0 & C_{55} & 0 \\ 0 & 0 & 0 & 0 & 0 & C_{66} \end{bmatrix}$$

$$C_{11} \approx M \left[1 - \delta_{N1} - (1 - 2g)^2 \delta_{N2} \right],$$

$$C_{12} \approx \lambda (1 - \delta_{N1} - \delta_{N2}),$$

$$C_{13} \approx \lambda [1 - \delta_{N1} - (1 - 2g) \delta_{N2}],$$

$$C_{22} \approx M \left[1 - (1 - 2g)^2 \delta_{N1} - \delta_{N2} \right],$$

$$C_{23} \approx \lambda [1 - (1 - 2g) \delta_{N1} - \delta_{N2}],$$

$$C_{33} \approx M \left[1 - (1 - 2g)^2 \delta_{N1} - (1 - 2g)^2 \delta_{N2} \right],$$

$$C_{44} = \mu (1 - \delta_{T2}),$$

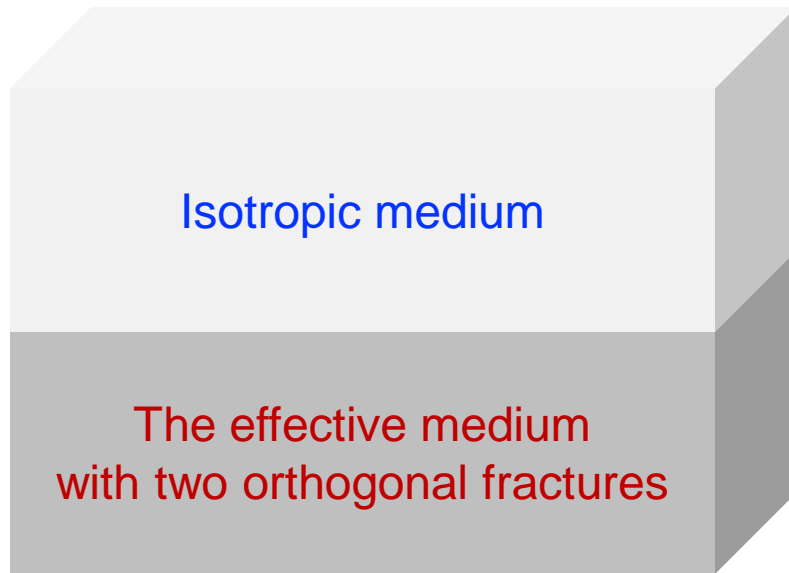
$$C_{55} = \mu (1 - \delta_{T1}),$$

$$C_{66} \approx \mu (1 - \delta_{T1} - \delta_{T2}).$$



Theory and method: PP- and PS-wave reflection coefficients

An interface separating an isotropic medium and a medium with two orthogonal fractures



$$R_{PP}(\theta, \phi) = R_{PP}^{\text{iso}}(\theta) + R_{PP}^{\text{ani}}(\theta, \phi),$$

$$R_{PS}(\theta, \phi) = R_{PS}^{\text{iso}}(\theta) + R_{PS}^{\text{ani}}(\theta, \phi)$$

Isotropic part

$$R_{PP}^{\text{iso}}(\theta) = a_{\rho}(\theta) \frac{\Delta\rho}{\rho} + a_M(\theta) \frac{\Delta M}{M} + a_{\mu}(\theta) \frac{\Delta\mu}{\mu},$$

$$R_{PS}^{\text{iso}}(\theta) = b_{\mu}(\theta) \frac{\Delta\mu}{\mu} + b_{\rho}(\theta) \frac{\Delta\rho}{\rho}$$

Anisotropic part

$$R_{PP}^{\text{ani}}(\theta, \phi) = a_{N1}(\theta, \phi)\delta_{N1} + a_{N2}(\theta, \phi)\delta_{N2} \\ + a_{T1}(\theta, \phi)\delta_{T1} + a_{T2}(\theta, \phi)\delta_{T2},$$

$$R_{PS}^{\text{ani}}(\theta, \phi) = b_{N1}(\theta, \phi)\delta_{N1} + b_{N2}(\theta, \phi)\delta_{N2} \\ + b_{T1}(\theta, \phi)\delta_{T1} + b_{T2}(\theta, \phi)\delta_{T2},$$



Theory and method: PP- and PS-wave reflection coefficients

Assumptions for simplifying two sets of fracture weaknesses

1) The same infilling fluids

$$K_f = K_{f1} = K_{f2}$$

2) The approximately equal aspect ratio

$$\chi \approx \chi_1 \approx \chi_2$$

$$\delta_{N1} = \frac{4e_1}{3g(1-g) \left[1 + \frac{1}{\pi g(1-g)} \frac{K_{f1}}{\mu\chi_1} \right]}$$

$$\delta_{T1} = \frac{16e_1}{3(3-2g)}$$

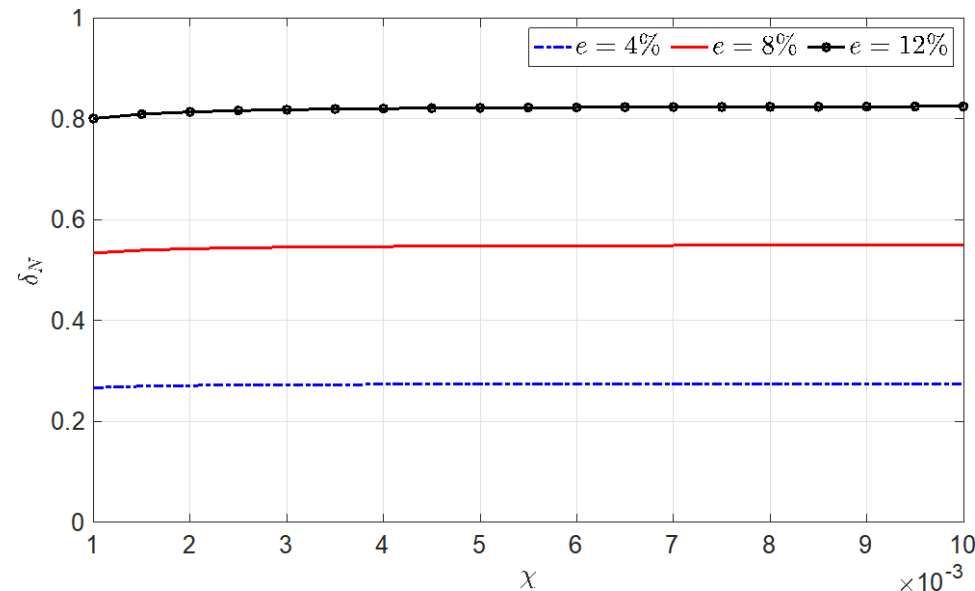
$$\delta_{N2} = \frac{4e_2}{3g(1-g) \left[1 + \frac{1}{\pi g(1-g)} \frac{K_{f2}}{\mu\chi_2} \right]}$$

$$\delta_{T2} = \frac{16e_2}{3(3-2g)}$$

$$\delta_{N1} = \gamma \delta_{N2},$$

$$\delta_{T1} = \gamma \delta_{T2},$$

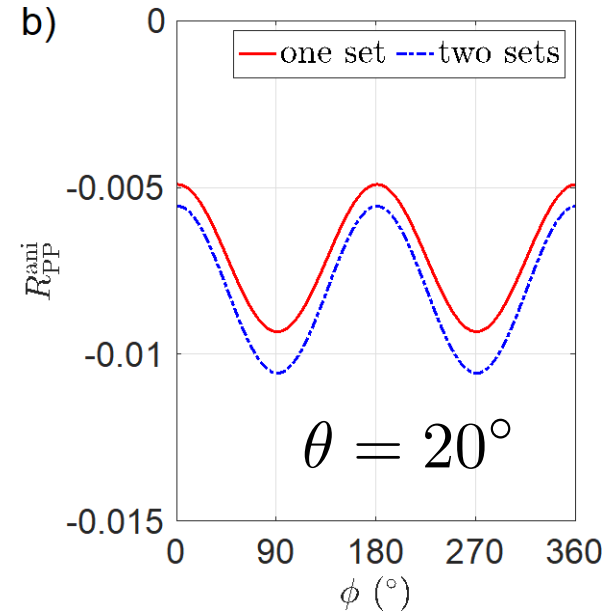
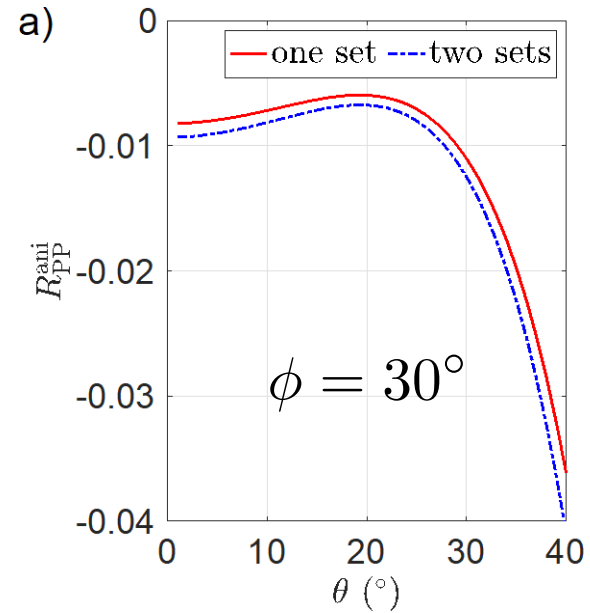
$$\gamma = \frac{e_1}{e_2}$$



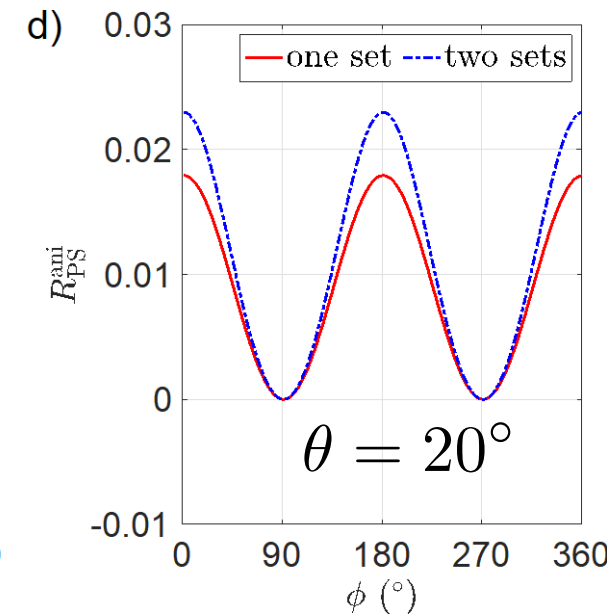
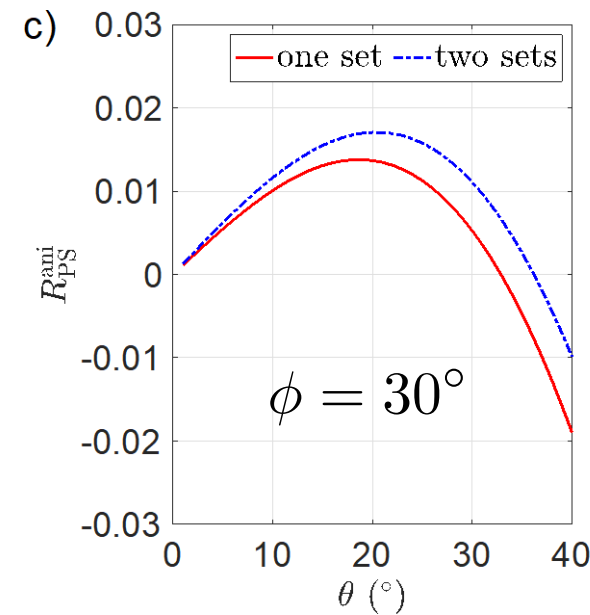


Theory and method: PP- and PS-wave reflection coefficients

	δ_{N1}	δ_{T1}	δ_{N2}	δ_{T2}
Sand	0	0	0	0
Shale (one set)	0	0	0.7849	0.3183
Shale (two sets)	0.1047	0.0424	0.7849	0.3183



PP-wave



PS-wave



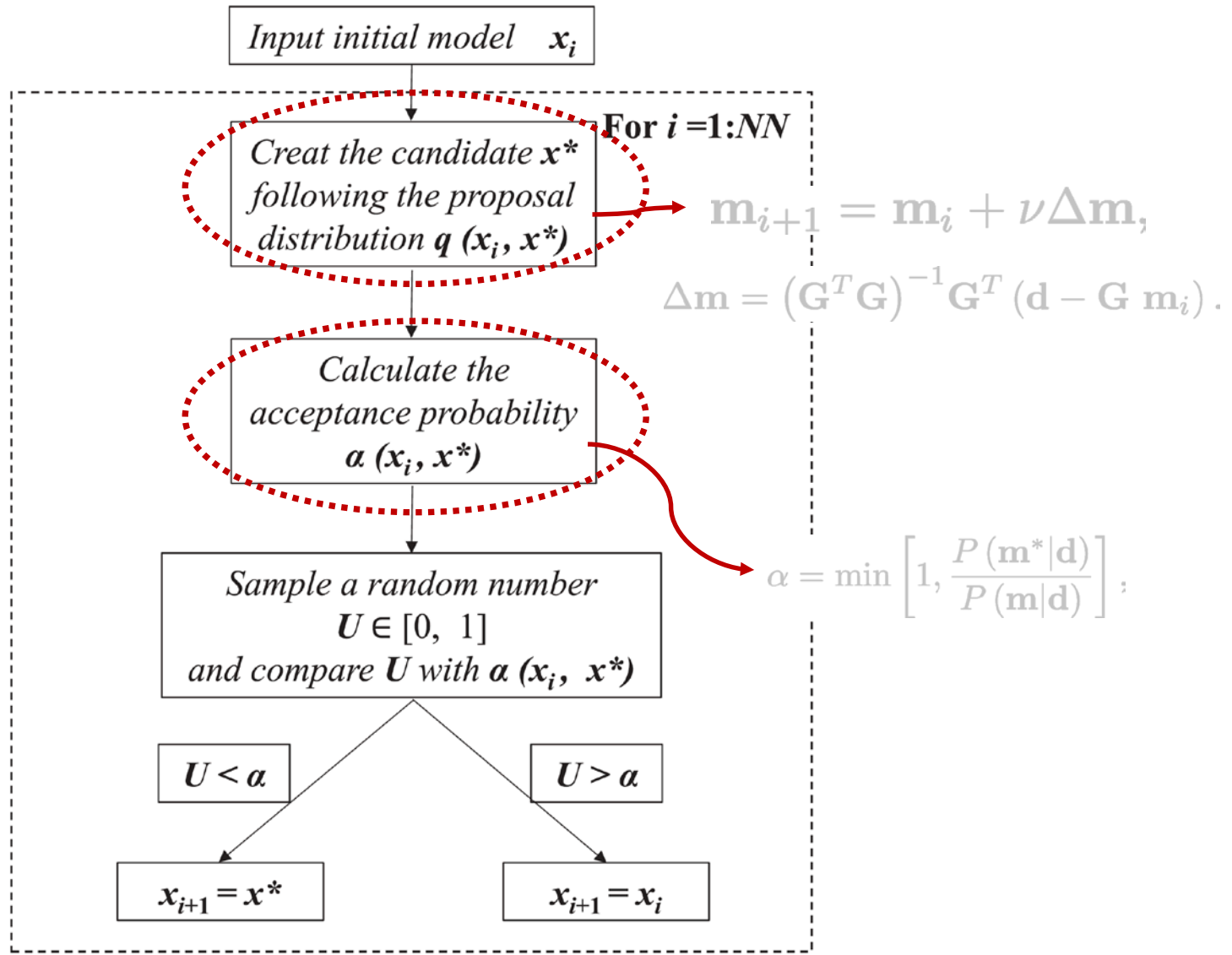
Input datasets:

Differences in azimuthal reflection coefficients variation with incidence angle

$$\Delta R_{PP}(\theta, \phi_1, \phi_k) = \mathcal{P}_N(\theta, \phi_1, \phi_k) \delta_{N2} + \mathcal{P}_T(\theta, \phi_1, \phi_k) \delta_{T2},$$

$$\Delta R_{PS}(\theta, \phi_1, \phi_k) = \mathcal{Q}_N(\theta, \phi_1, \phi_k) \delta_{N2} + \mathcal{Q}_T(\theta, \phi_1, \phi_k) \delta_{T2},$$

Workflow





Posterior PDF

$$P(\mathbf{m}|\mathbf{d}) \propto P(\mathbf{d}|\mathbf{m})P(\mathbf{m})$$

$$\propto \exp\left(-\frac{(\mathbf{d} - \mathbf{Gm})^T (\mathbf{d} - \mathbf{Gm})}{2\sigma_{\text{noise}}^2}\right) \exp\left(-\frac{(\mathbf{m} - m_a)^T (\mathbf{m} - m_a)}{2\sigma_m^2}\right)$$

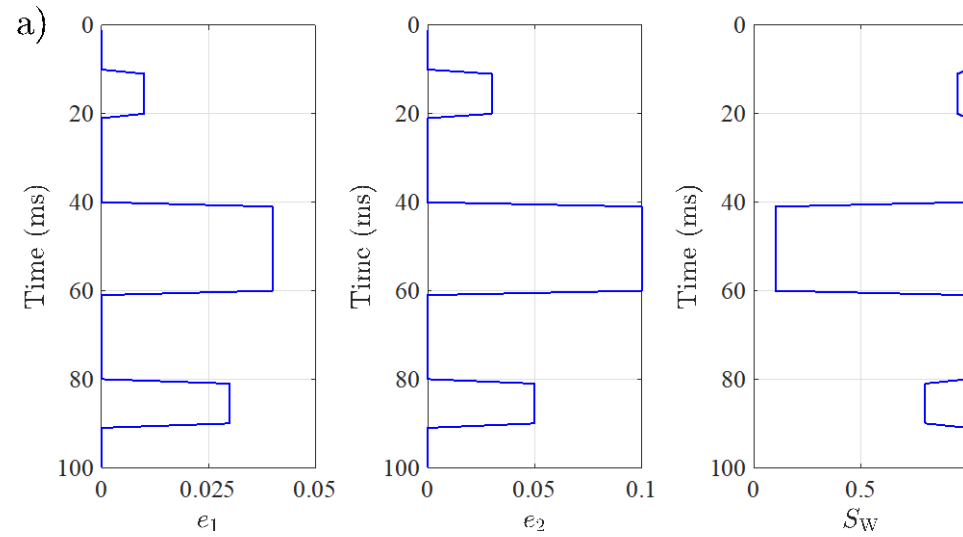
Gaussian random noise

Gaussian priori

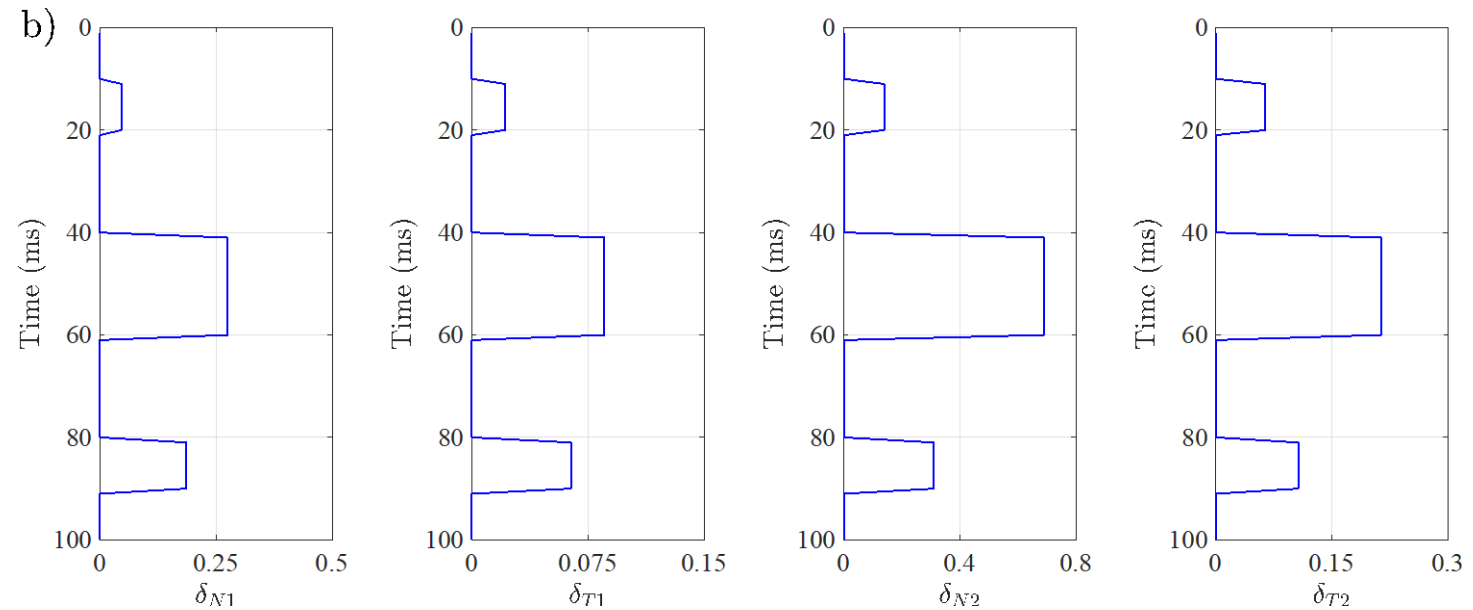


Numerical examples

Curves of two fracture density and water saturation

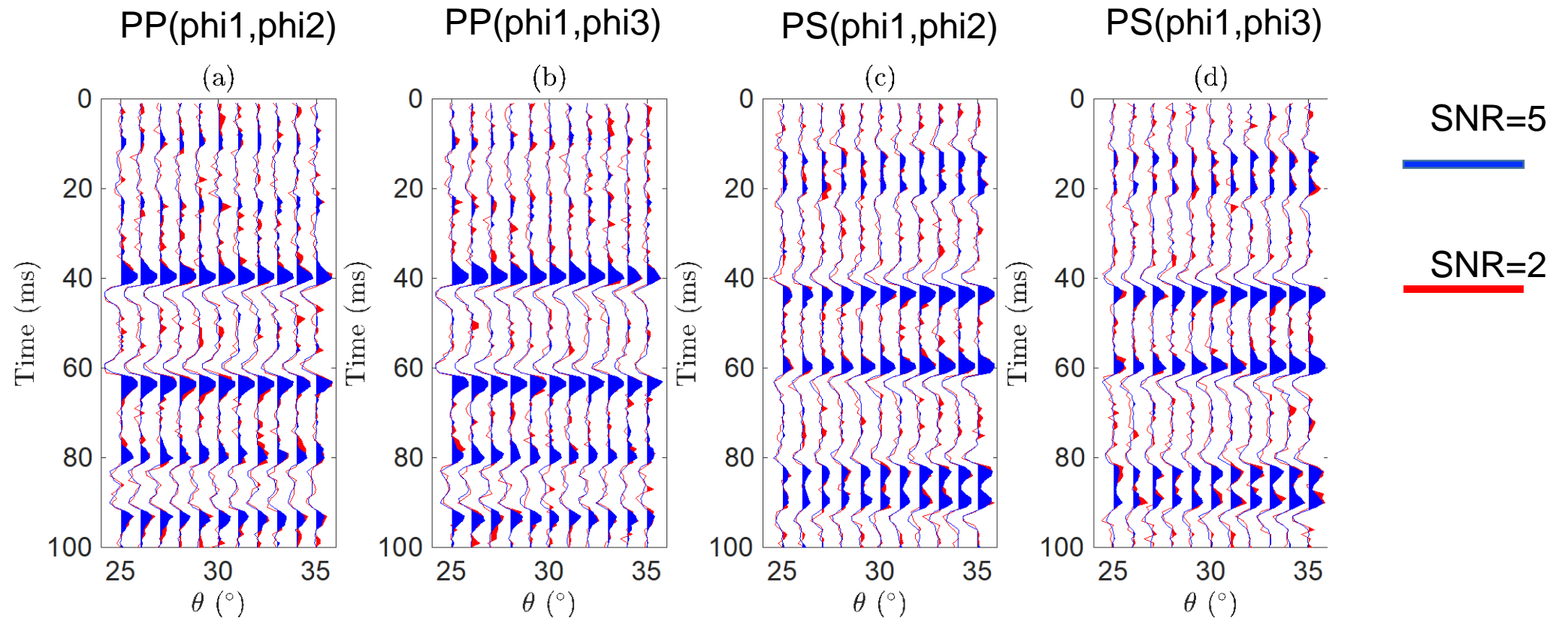


Calculated fracture weaknesses





Noisy synthetic seismic data

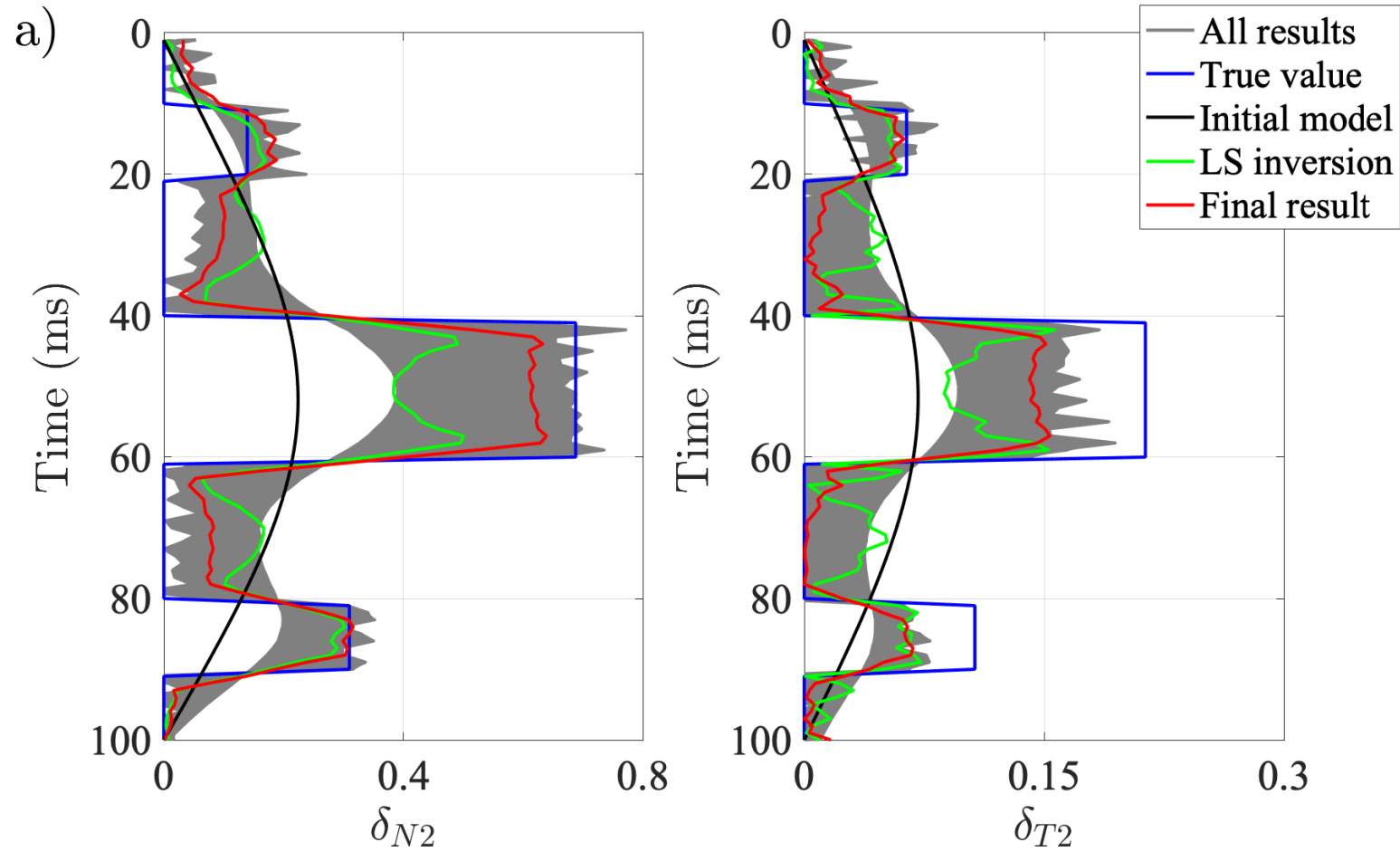


$$\phi_1 = 0^{\circ}, \phi_2 = 45^{\circ}, \phi_3 = 90^{\circ}$$



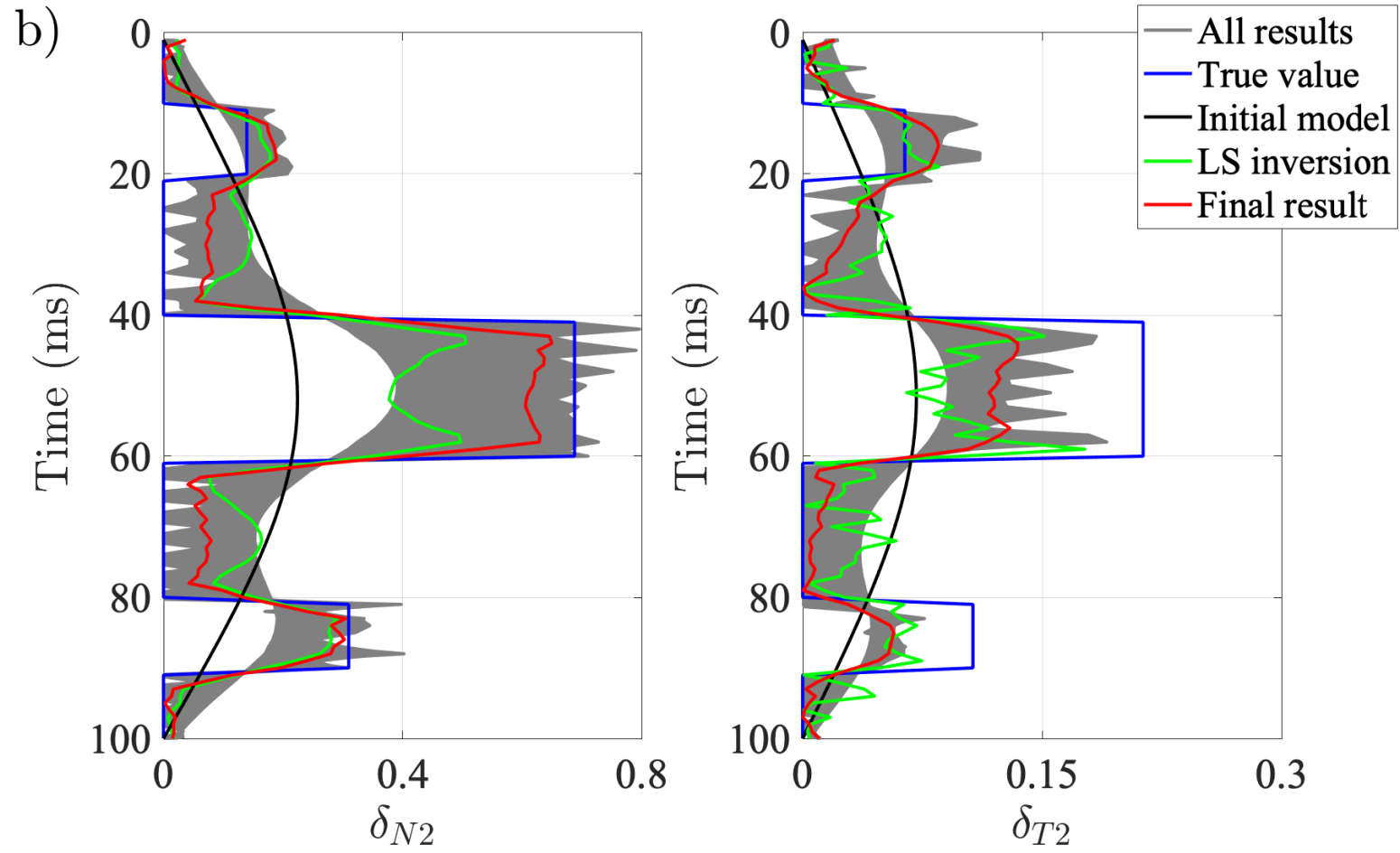
Numerical examples

SNR=5



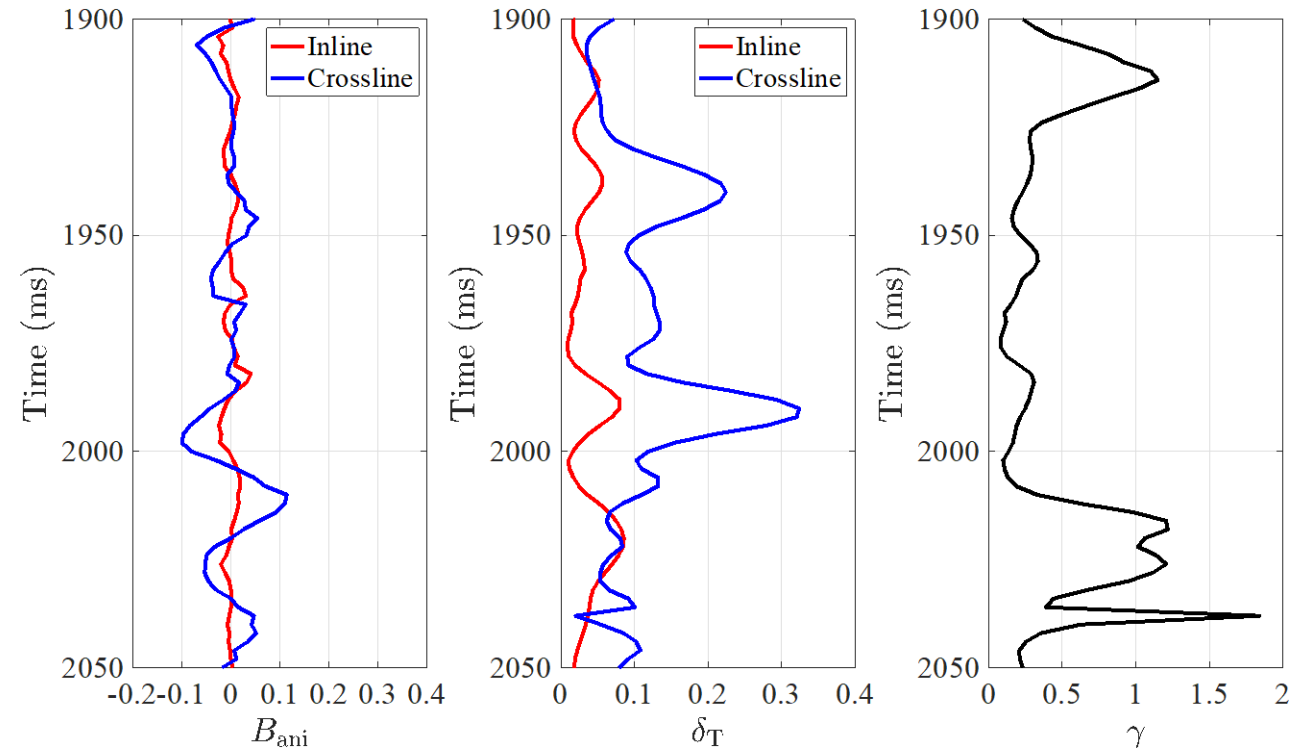
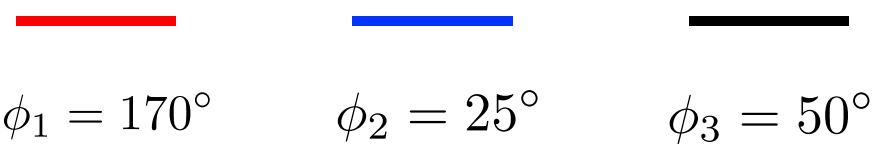
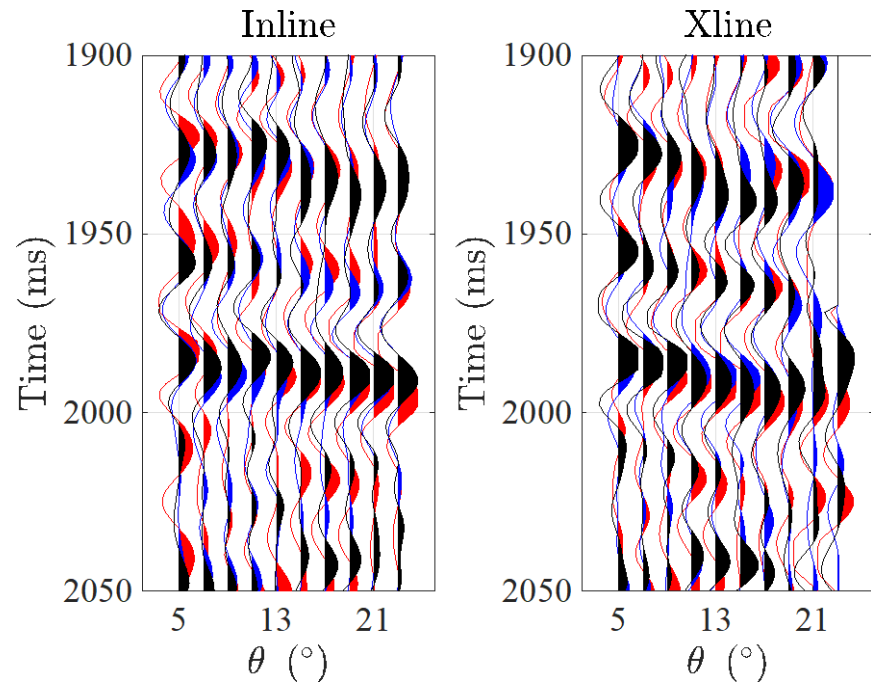


SNR=2





Real datasets



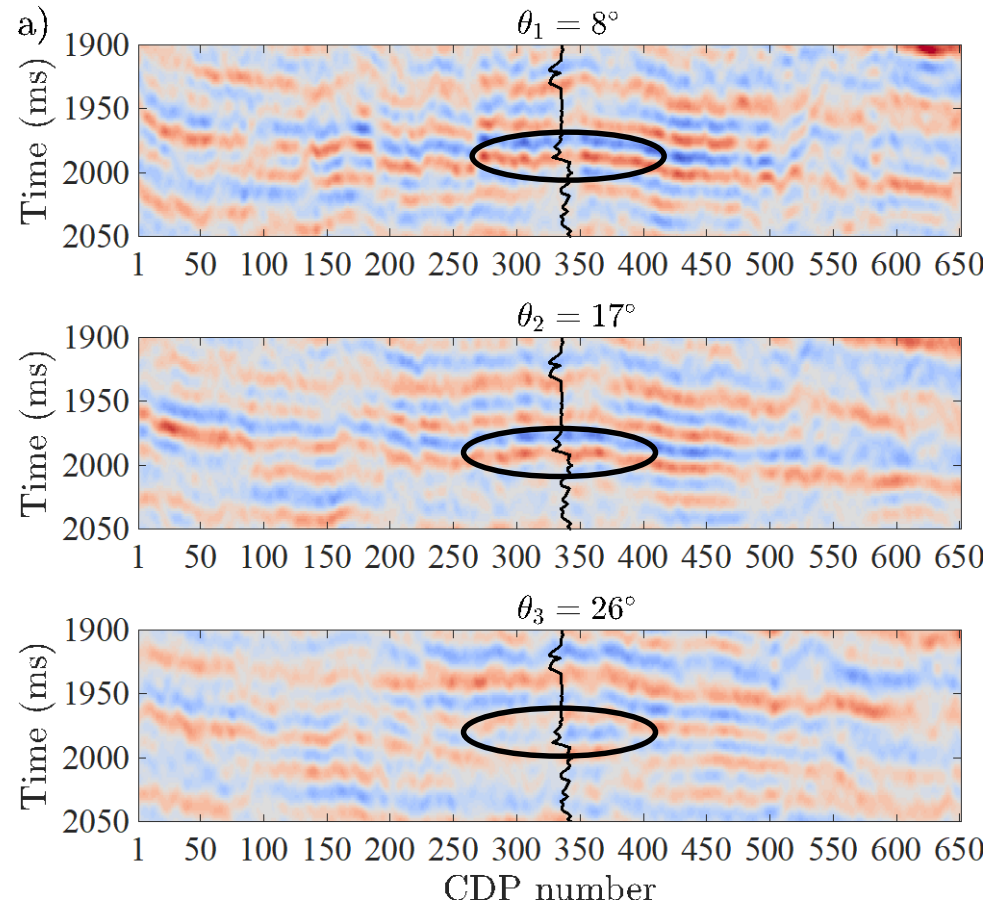
$$B_{\text{ani}} \approx -g(1-2g)\delta_N + g\delta_T$$

$$\delta_N \approx \frac{3-2g}{4g(1-g)}\delta_T$$

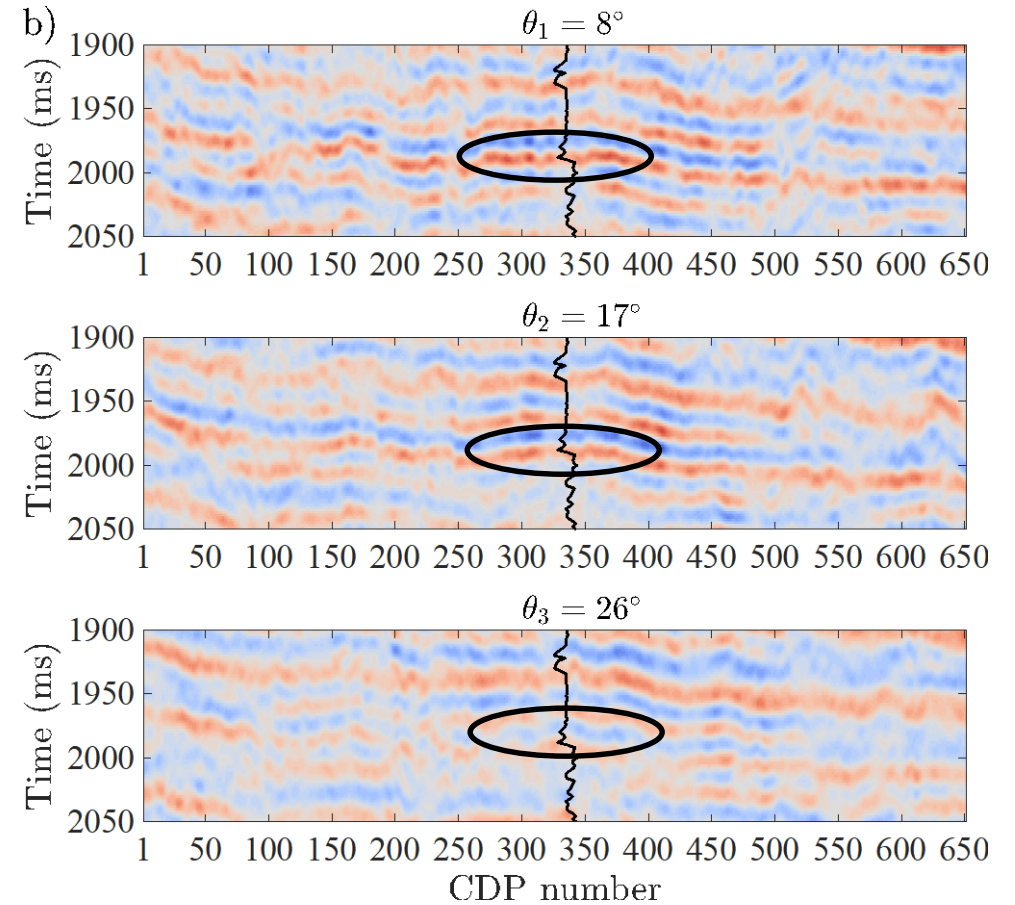


Numerical examples

$$\phi_1 = 170^\circ$$



$$\phi_2 = 25^\circ$$

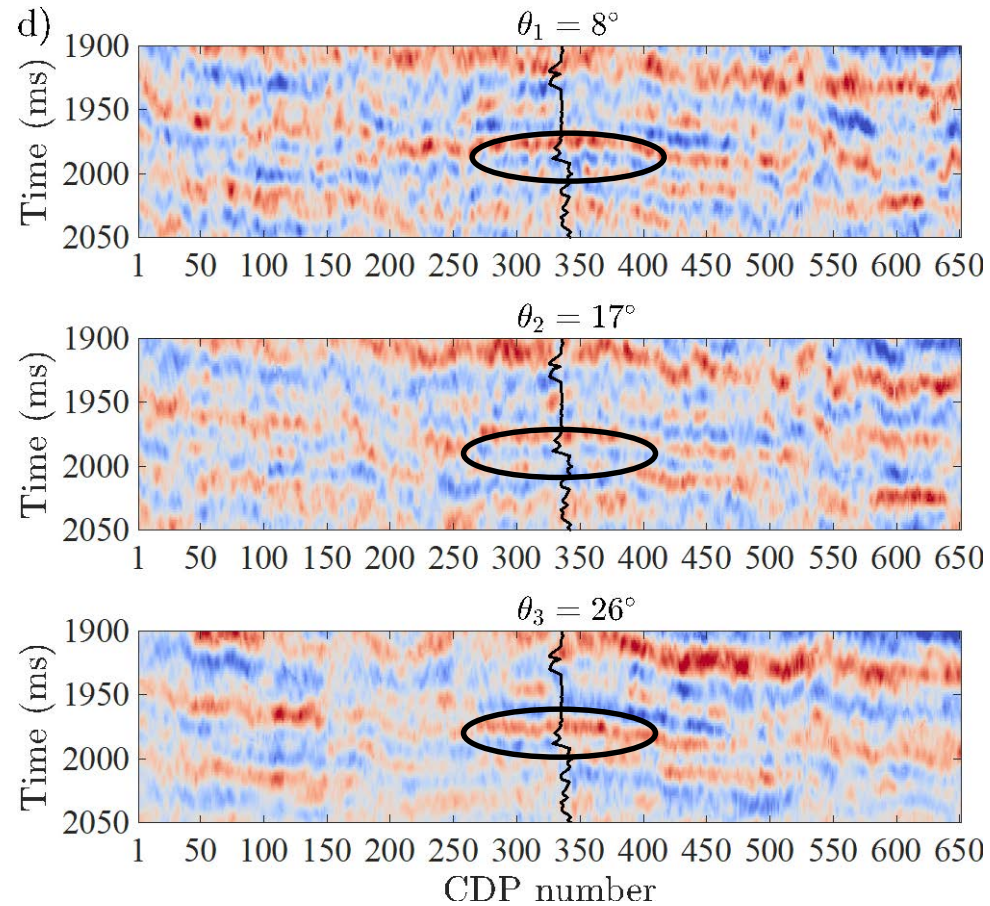


PP-wave stacked data of small, middle and large angles

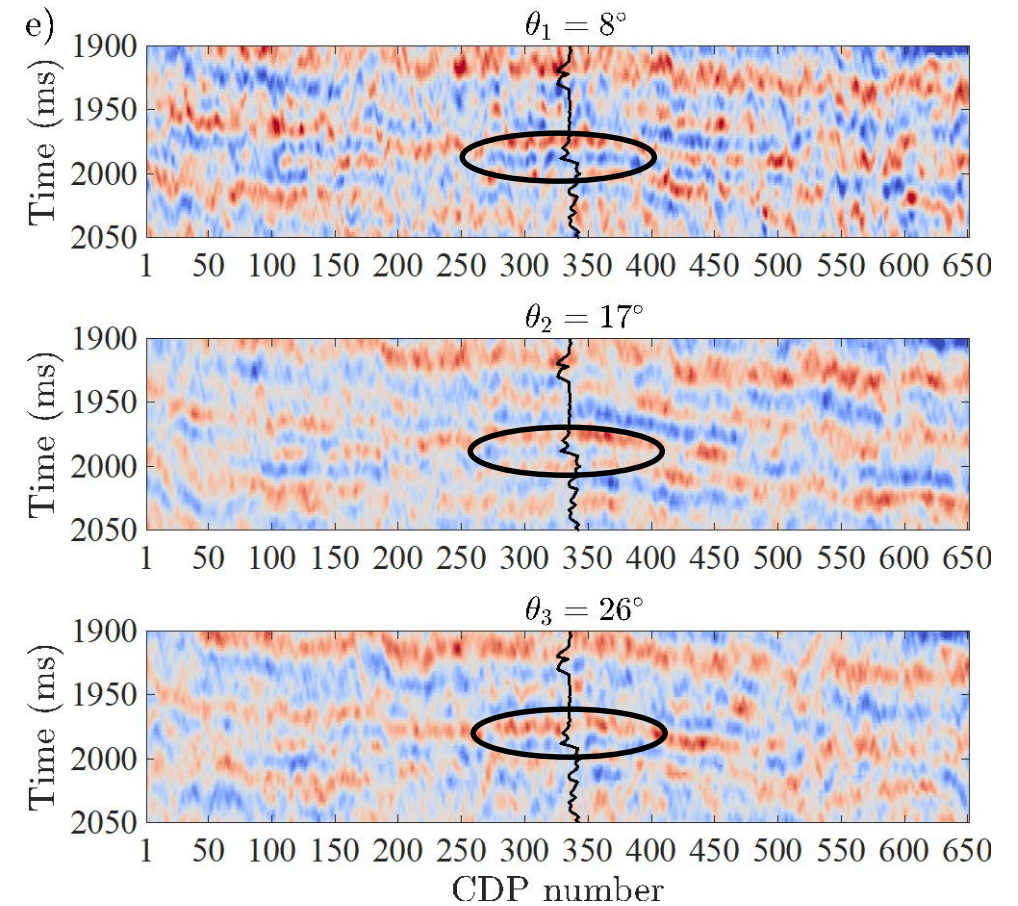


Numerical examples

$$\phi_1 = 170^\circ$$



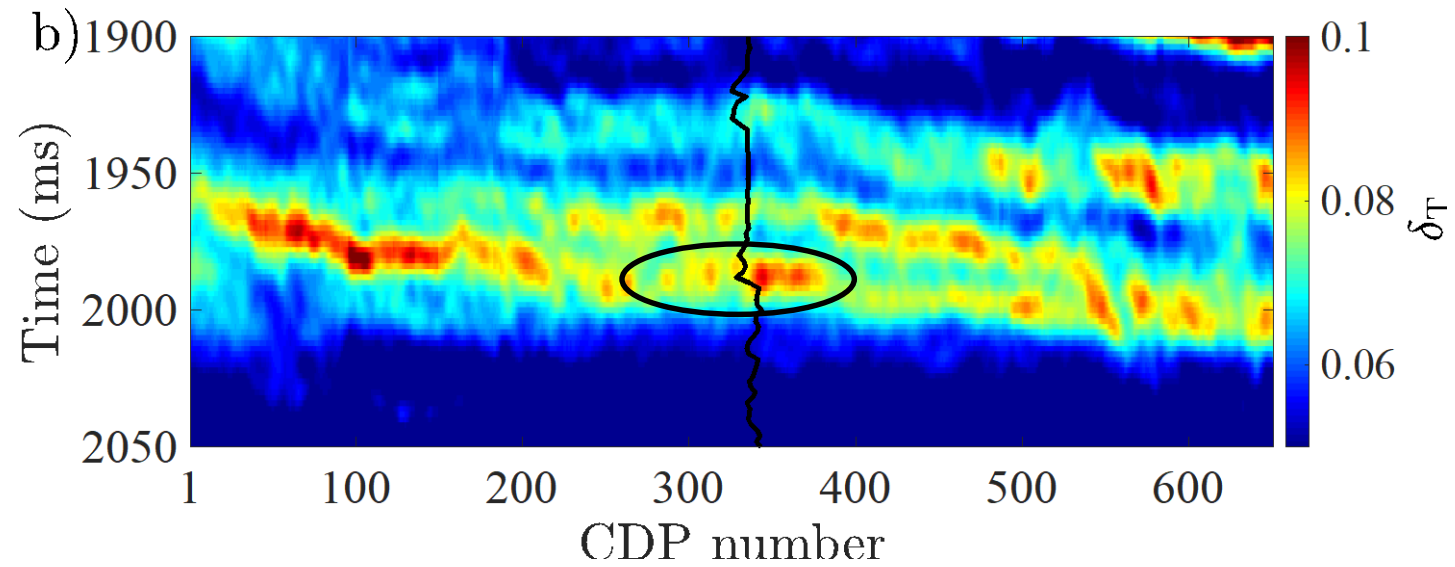
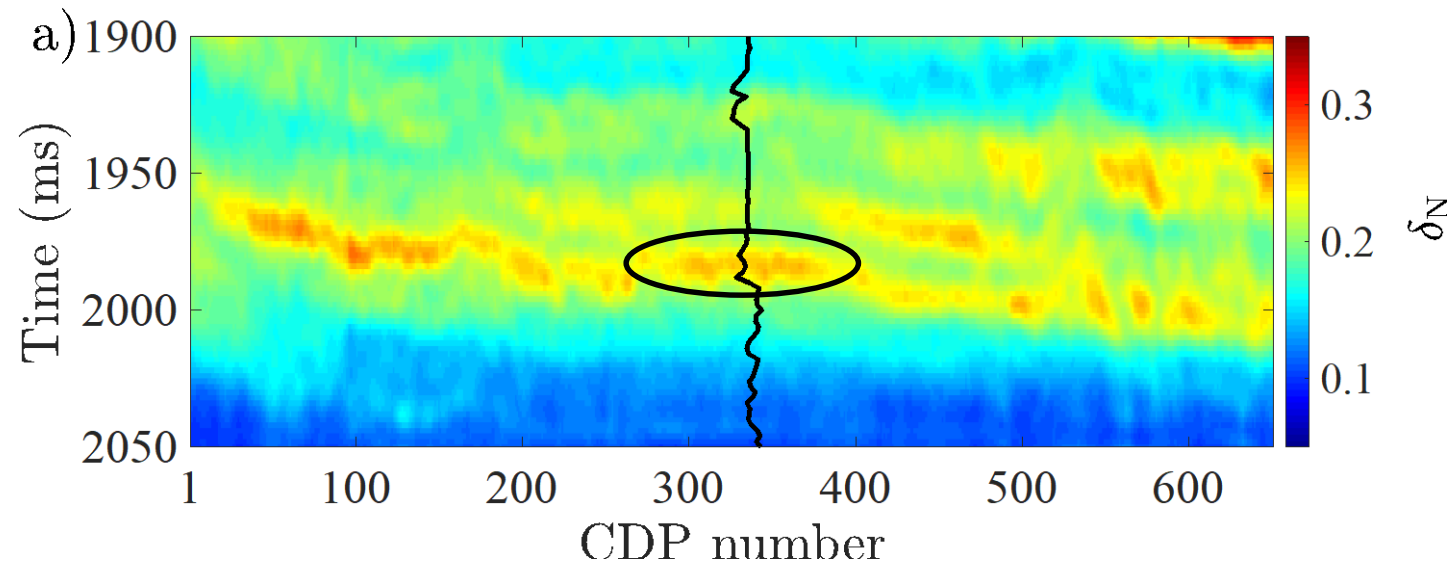
$$\phi_2 = 25^\circ$$



PS-wave stacked data of small, middle and large angles



Numerical examples





Conclusions

- We derive linearized PP- and PS-wave reflection coefficients in terms of two sets of fracture weaknesses to model how fracture connectivity affects reflection coefficients;
- We establish an inversion approach combining LS algorithm and Bayesian MCMC method to estimate fracture weaknesses from azimuthal seismic amplitudes;
- Tests on noisy synthetic data and real data reveal that the proposed approach is stable and robust and can provide the possibility to determine how fractures distribute and whether fractures are connected.



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